

Implementation and Validation of a Novel Representation of the Gravitational Effect of Ocean Tides

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Ocean tides

- Tides remain a **crucial contributor** to a variety of geodetic applications, including satellite altimetry and gravimetry.
 - In recent years, **significant advances** have been made in ocean tide models in terms of model accuracies as well as in model abilities to derive more tidal constituents.
 - Empirical, data-driven, tide models produce the highest level of accuracy largely thanks to satellite altimetry.
 - Purely hydrodynamic models are not restricted by aliasing or noise related constraints allowing them to estimate a wider range of tidal constituents.
- **Data formats allowing to flexibly apply different tide models in space-geodetic analysis are important!**

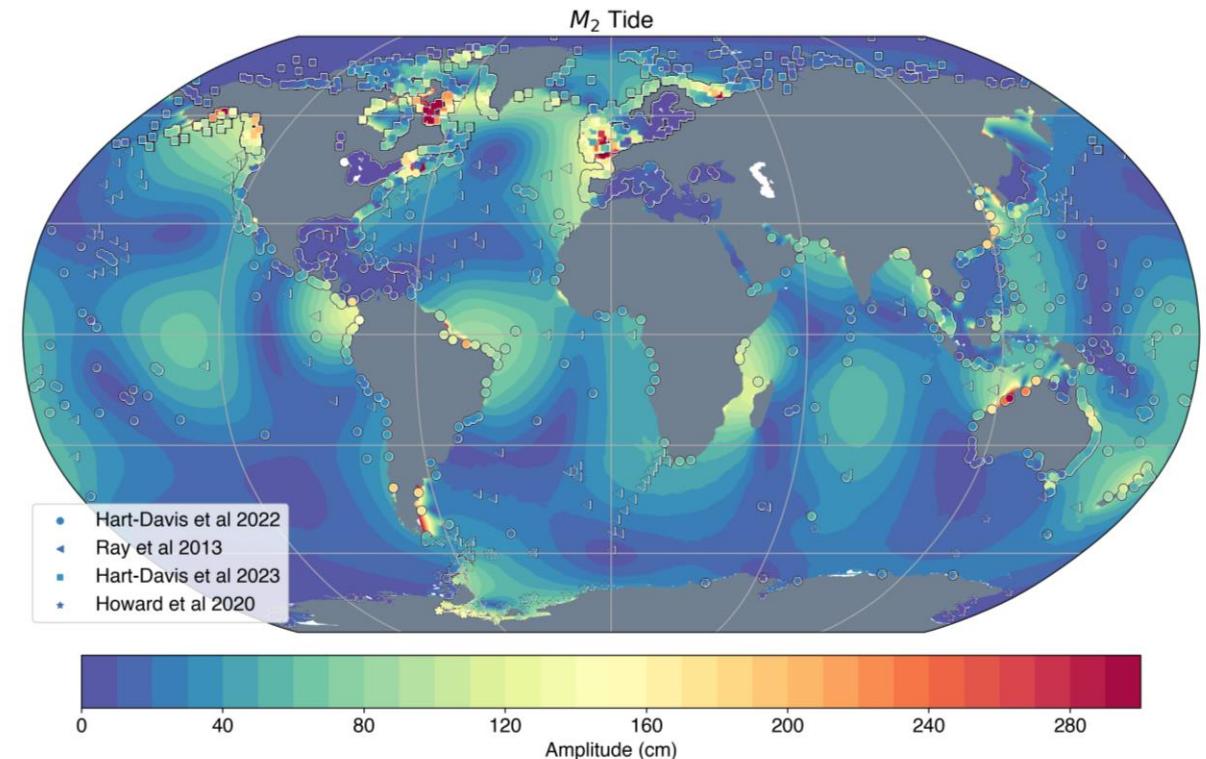


Fig. The M_2 tide from EOT20 tide model (Hart-Davis et al 2021) overlaid with data from available in-situ tidal constituent databases.

Parameterisation of the gravitational effect of ocean tides

IERS Conventions 2010 (Petit and Luzum, 2010, Sect. 6.3):

- Ocean tides are provided in the form of spherical harmonics for tide heights
- Synthesis with period-dependent phase biases (Doodson-Warburg convention; IERS Conv. 2010, Tab. 6.6)
- Tidal admittance via „hardcoded“ table (adapted to the old FES2004 model; IERS Conv. 2010, Tab. 6.7)

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- Various sources for mis-implementation
 - Implementational effort for any new model
 - Inconsistencies due to model-dependent ambiguous tide definitions (e.g. S1: 164.556 or 164.555)

Alternative approach (Mayer-Gürr et al., 2023) based on the standard **ICGEM format**:

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- Ocean tides are provided in the form of spherical harmonics for the gravitational effect
 - No phase biases
 - Tidal admittance and Doodson multipliers provided in matrices
 - Model-independent and unambiguous implementation
 - Any tidal admittance method can be used
 - **Implemented and tested in DOGS-OC**

Implementation

Ocean tide synthesis at time t

$$\begin{bmatrix} c_{nm} \\ s_{nm} \end{bmatrix}(t) = \sum_k f_k^{\cos}(t) \begin{bmatrix} c_{nm} \\ s_{nm} \end{bmatrix}_k^{\cos} + f_k^{\sin}(t) \begin{bmatrix} c_{nm} \\ s_{nm} \end{bmatrix}_k^{\sin}$$

Temporal changing factors

$$f_k^{\cos}(t) = \sum_f A_{k,f} \cos \theta_f(t),$$

$$f_k^{\sin}(t) = \sum_f A_{k,f} \sin \theta_f(t)$$

Admittance matrix

1.00000e+00	0.00000e+00	0.00000e+00	...
0.00000e+00	1.00000e+00	-1.12052e-01	...
0.00000e+00	0.00000e+00	-1.48522e-03	...
...			

All tidal lines are treated in the same way

Flexible: different interpolation schemes,
adding non TGP tides, equilibrium tides, resonances...

Fast

Source:

Mayer-Gürr et al. (2023): *Exploiting the full potential of ocean tide models for space geodetic techniques*. EGU General Assembly 2023, DOI 10.5194/egusphere-egu23-13235

Phase arguments for all tidal lines

$$\theta_f(t) = \sum_{i=1}^6 D_{f,i} \beta_i(t) \quad \leftarrow \text{6 Doodson arguments}$$

Matrix with Doodson multipliers

0	0	0	0	1	0
0	0	0	0	2	0
0	0	0	2	1	0
0	0	1	0	-1	-1
0	0	1	0	0	-1
...					

Do not care about

- Darwin names / Doodson codes
- Doodson-Warburg phase shifts

Comparisons

▪ Solution setups:

- Models: **EOT11a** (Savcenko and Bosch, 2012), **EOT20** (Hart-Davis et al., 2020)
- Resolution: Up to degree/order 30 (low), 90 (middle), 180 (high)
- Tidal admittance: none (only main tides), 63 secondary tides (IERS Conv. 2010), 335 secondary tides

▪ Satellites:

- LAGEOS-1 (weekly arcs)
- LARES (weekly arcs)
- Jason-3 (3.5-day arcs)

▪ Parameterisation:

- solar radiation pressure scaling factor
 - atmospheric drag pwl scaling factor (LARES, Jason-3)
 - along-track pwl empirical forces (LAGEOS-1)
 - cos/sin transversal/normal empirical forces
-
- station coordinates and Earth rotation parameters fixed
 - biases according to ILRS Data Handling File

Note:

- Gravitational effect of atmospheric tides not applied
- Ocean (EOT11a/Scherneck¹) and atmospheric (Ray and Ponte, 2003) loading unchanged

¹<http://holt.oso.chalmers.se/loading/>

Results: Impact of tidal admittance and resolution

- Applying tidal admittance reduces the arc RMS by multiple millimetres:

EOT11a with	<i>18 main 0 secondary</i>	<i>18 main 63 secondary</i>	<i>18 main 335 secondary</i>
(IERS 2010)			

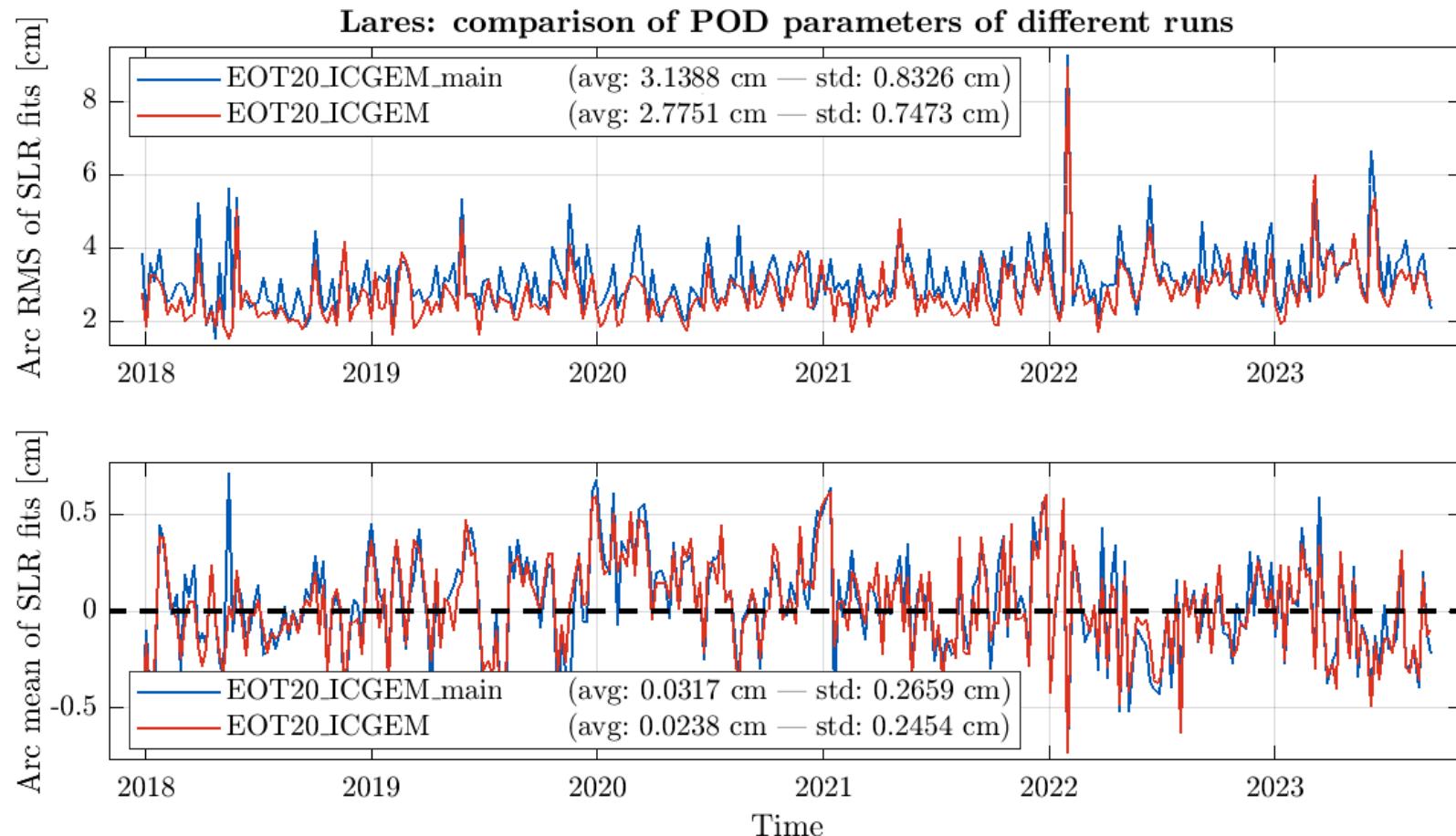
Satellite	altitude	arc rms of SLR fits [cm]		
LAGEOS-1	5850 km	1.2108	1.1931	1.1928
LARES	1450 km	3.1290	2.7630	2.7546
Jason-3	1336 km	2.5687	2.4275	2.4263

- For the satellites used in this study, the maximum degree/order (30, 90, or 180) has no significant impact on the results
- Small differences between the results for EOT11a and EOT20
 - EOT20 long-period tides could contain non-tidal loading effects which are not present in EOT11a
 - We expect that these cause the changes in SLR RMS fits

EOT20 with	<i>19 main 0 secondary</i>	<i>19 main 335 secondary</i>
(IERS 2010)		
Satellite	arc rms of SLR fits [cm]	
LAGEOS-1	1.2111	1.1939
LARES	3.1388	2.7751
Jason-3	2.5658	2.4254

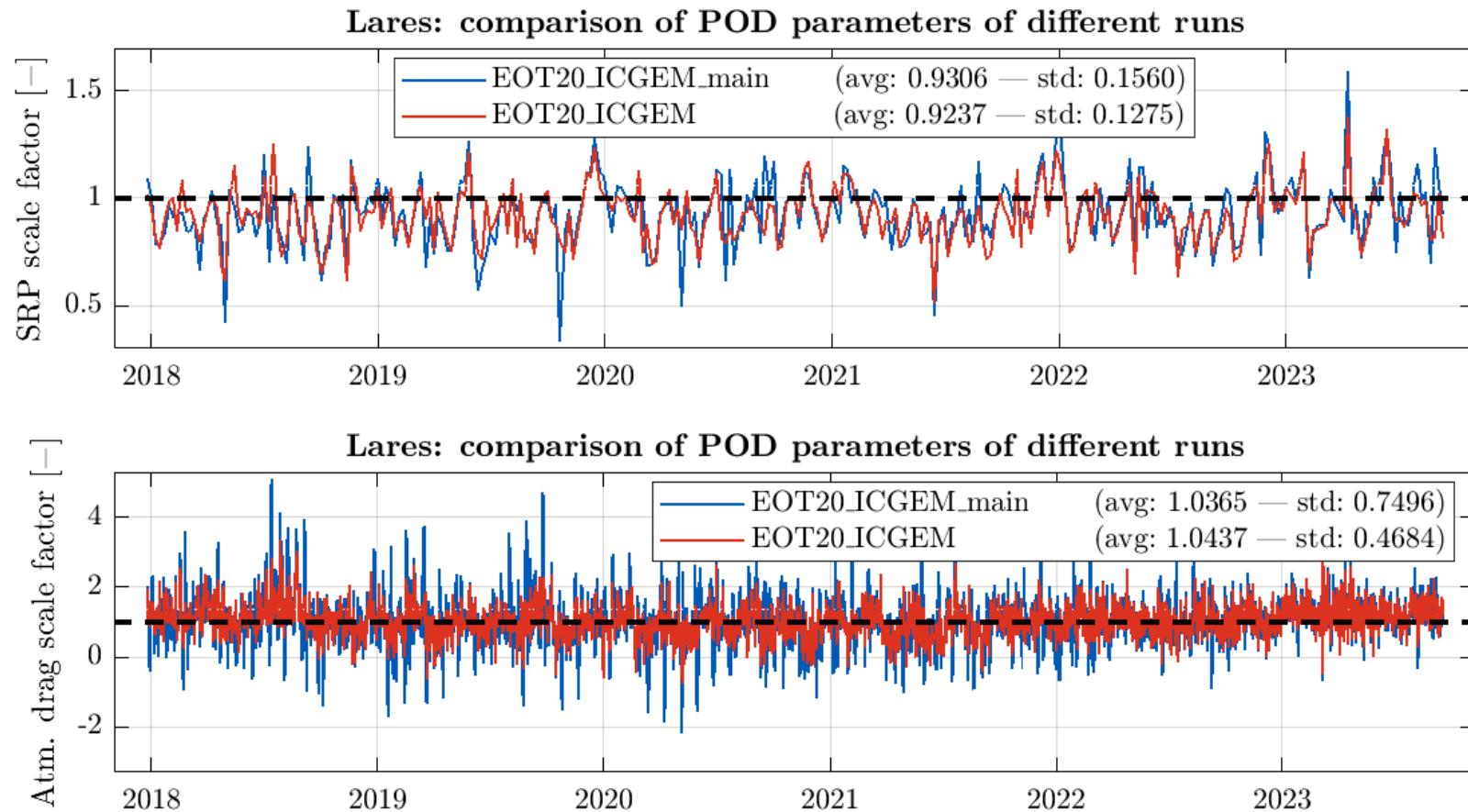
LARES: RMS of SLR fits

- RMS of SLR fits is reduced by several millimetres



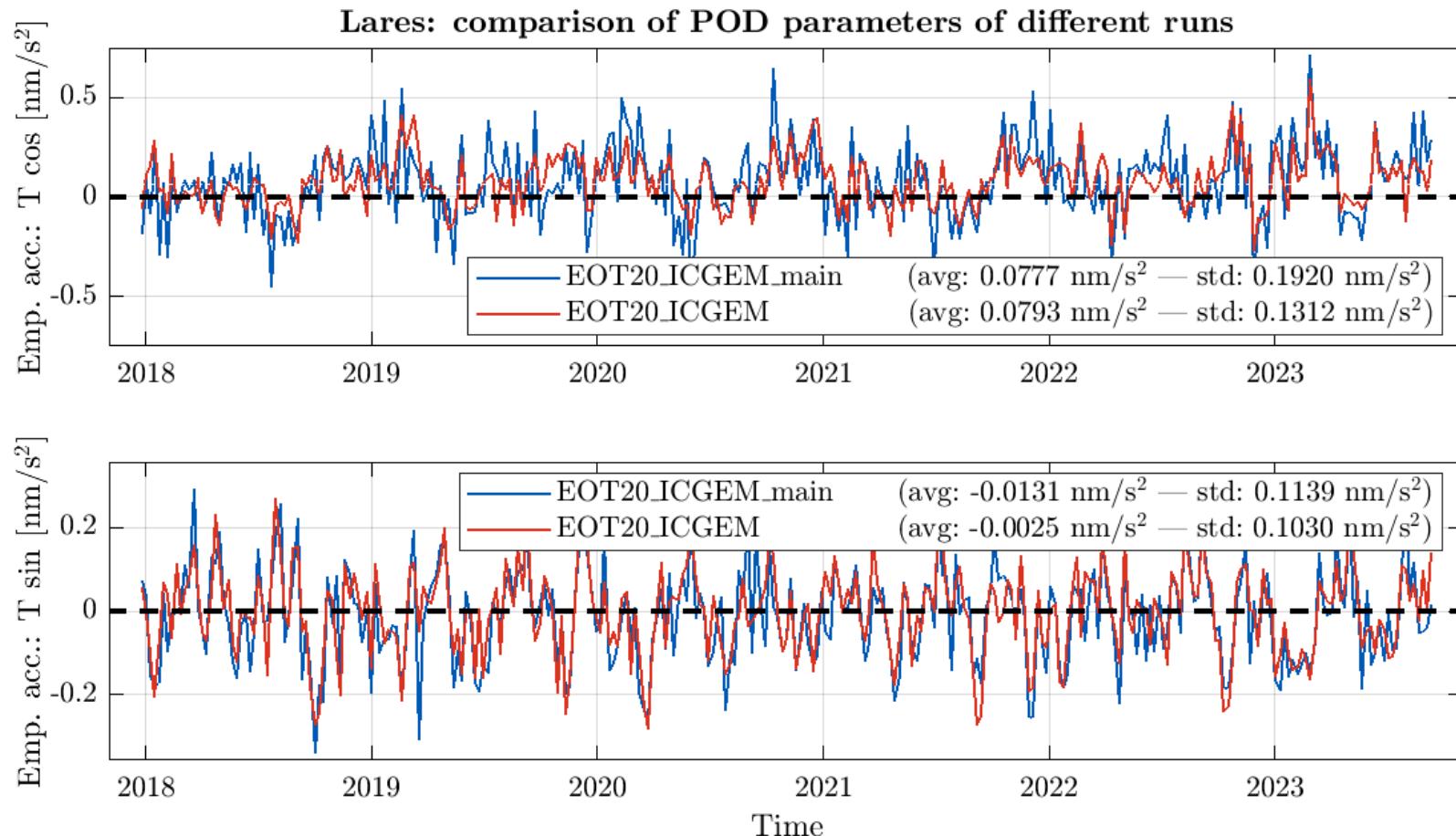
LARES: Solar radiation pressure and atmospheric drag scale factors

- Scatter of solar radiation pressure (SRP) and atmospheric drag scale factors is reduced



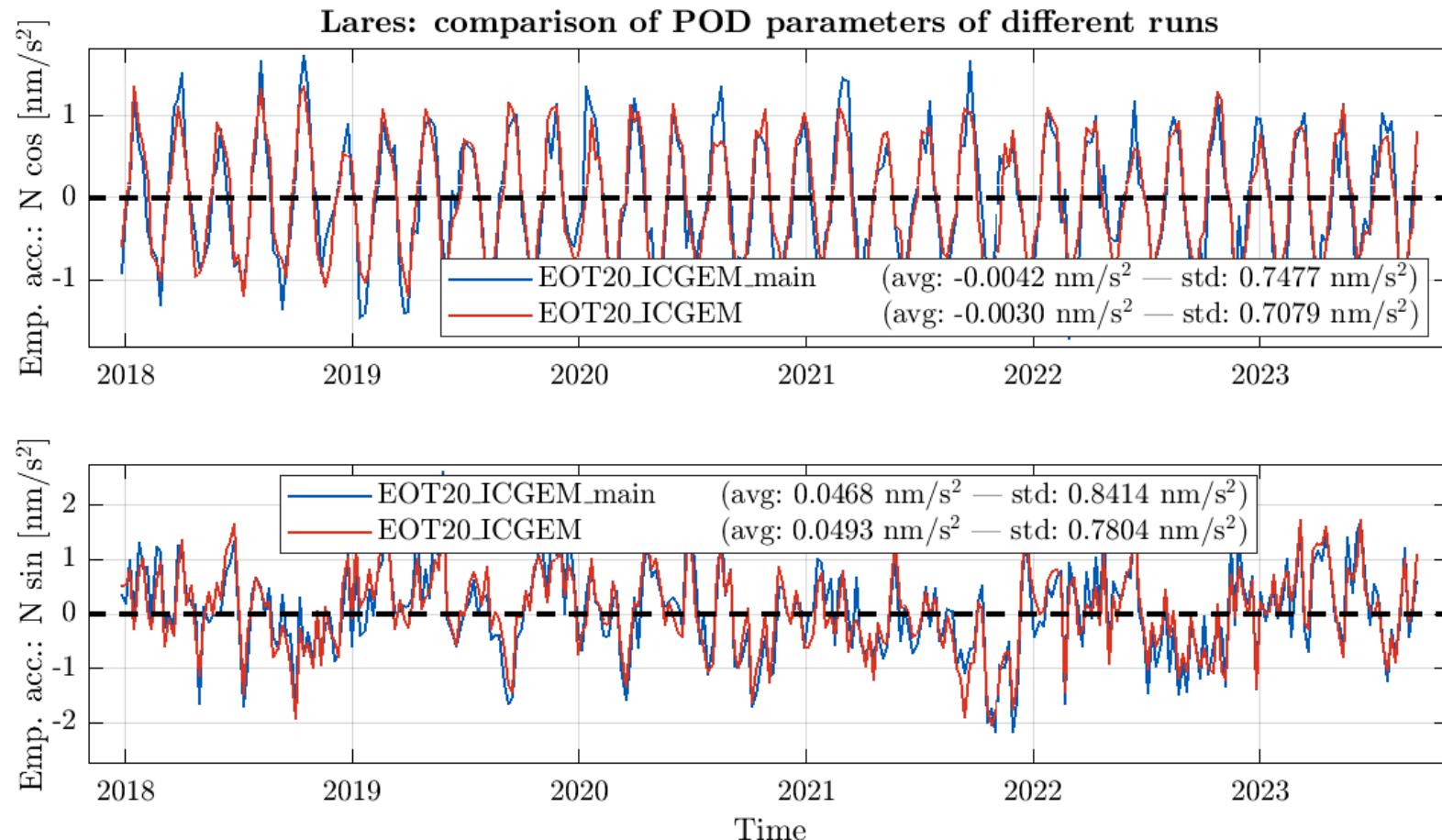
LARES: Empirical accelerations (transverse)

- Less scatter especially for the cosine component



LARES: Empirical acceleration (normal)

- Reduction in scatter for both cosine and sine component



Results: Impact of tidal admittance and resolution (EOT11a, EOT20)

<i>EOT20 with</i>	<i>19 main 0 secondary</i>	<i>19 main 335 secondary</i>	
LARES			
Arc RMS	[cm]	3.1388	2.7751
SRP scaling factor	[–]	0.9306 +/- 0.1560	0.9237 +/- 0.1275
Atmospheric drag scale factor	[–]	1.0365 +/- 0.7496	1.0437 +/- 0.4684
Emp. Acc. T cos/sin	[nm/s ²]	0.0777 0.0131 +/- 0.1920 +/- 0.1139	0.0793 -0.0025 +/- 0.1312 +/- 0.1030
Emp. Acc. N cos/sin	[nm/s ²]	-0.0042 0.0468 +/- 0.7477 +/- 0.8414	-0.0030 0.0493 +/- 0.7079 +/- 0.7804

Results: Impact of tidal admittance and resolution (EOT11a, EOT20)

<i>EOT20 with</i>		<i>19 main</i>		<i>19 main</i>	
		<i>0 secondary</i>		<i>335 secondary</i>	
Jason-3					
Arc RMS	[cm]	2.5658		2.4254	
SRP scaling factor	[–]	0.9876	+/- 0.0315	0.9879	+/- 0.0313
Atmospheric drag scale factor	[–]	1.2766	+/- 0.8956	1.2760	+/- 0.8878
Emp. Acc. T cos/sin	[nm/s ²]	0.0577	+/- 0.5236	0.0545	+/- 0.4954
		0.0115	+/- 0.3178	0.0234	+/- 0.3076
Emp. Acc. N cos/sin	[nm/s ²]	-0.0012	+/- 0.3638	-0.0170	+/- 0.3289
		-0.0473	+/- 0.5022	-0.0410	+/- 0.4507

Results: Impact of tidal admittance and resolution (EOT11a, EOT20)

<i>EOT20 with</i>	<i>19 main 0 secondary</i>	<i>19 main 335 secondary</i>	
LAGEOS-1			
Arc RMS	[cm]	1.2111	1.1939
SRP scaling factor	[–]	1.0349 +/- 0.0340	1.0339 +/- 0.0358
Atmospheric drag scale factor	[–]	–	–
Emp. Acc. T cos/sin	[nm/s ²]	0.0129 0.0001 +/- 0.0417 +/- 0.0327	0.0133 -0.0001 +/- 0.0409 +/- 0.0327
Emp. Acc. N cos/sin	[nm/s ²]	-0.0027 0.0195 +/- 0.1337 +/- 0.1224	-0.0046 0.0236 +/- 0.1323 +/- 0.1138

Conclusions

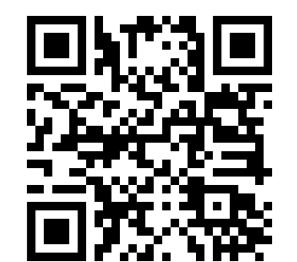
The results support using ocean tide models in ICGEM format with tidal admittance matrix:

- Model-independent implementation
- Less sources of ambiguity
- Flexible matrix-based tidal admittance
- Using more interpolated secondary tides
 - reduces the SLR orbit RMS
 - stabilises or reduces extrema of empirical forces estimates
 - stabilises solar radiation/atmospheric pressure scaling factors
- For classical SLR satellites, max. degree/order 30 seems sufficient
 - should be investigated for LEOs in lower orbits
- Reduction of orbit parameters absorbing unmodelled effects will allow for a **enhanced dynamic orbit estimation.**
- **Future work: test on other (combinations of) tide models.**

<https://ifg.tugraz.at/ocean-tides>

Ocean tides

- FES2014b
- EOT20, EOT11a
- TiME22
- ... Further models follow



Atmospheric tides

- TiME22
- ...

Reference implementations

- MATLAB, Python, Fortran

Skripts for converting ocean tide models

- from gridded NetCDF grids to spherical harmonics
- generating all necessary files
- based on GROOPS
- <https://github.com/groops-devs/groops>

References

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